

Land Level Change and Relative Sea Level Change

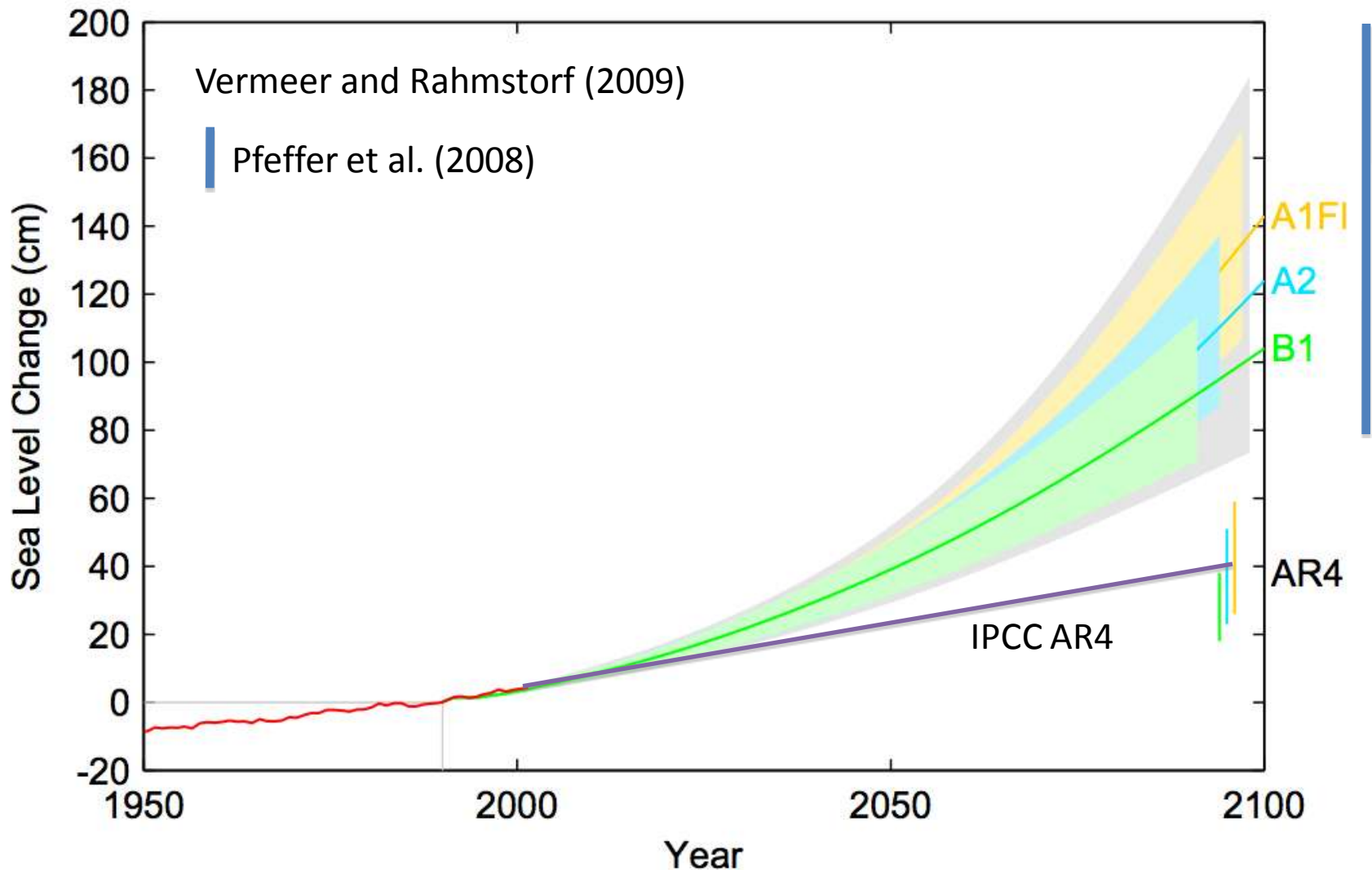
Jeff Freymueller
Geophysical Institute
University of Alaska Fairbanks



Take Home Messages

- Relative sea level is determined by both the level of the ocean and the level of the land
 - *The land surface in southern Alaska is moving faster than global sea level is presently changing.*
 - But rising sea level is likely to “catch up” eventually.
- We can measure these motions precisely
- We know the main causes of these motions
- We must extend models that explain horizontal motions to explain vertical motions.
 - Can we project motions decades into the future?

Predicted Global Average Sea Level



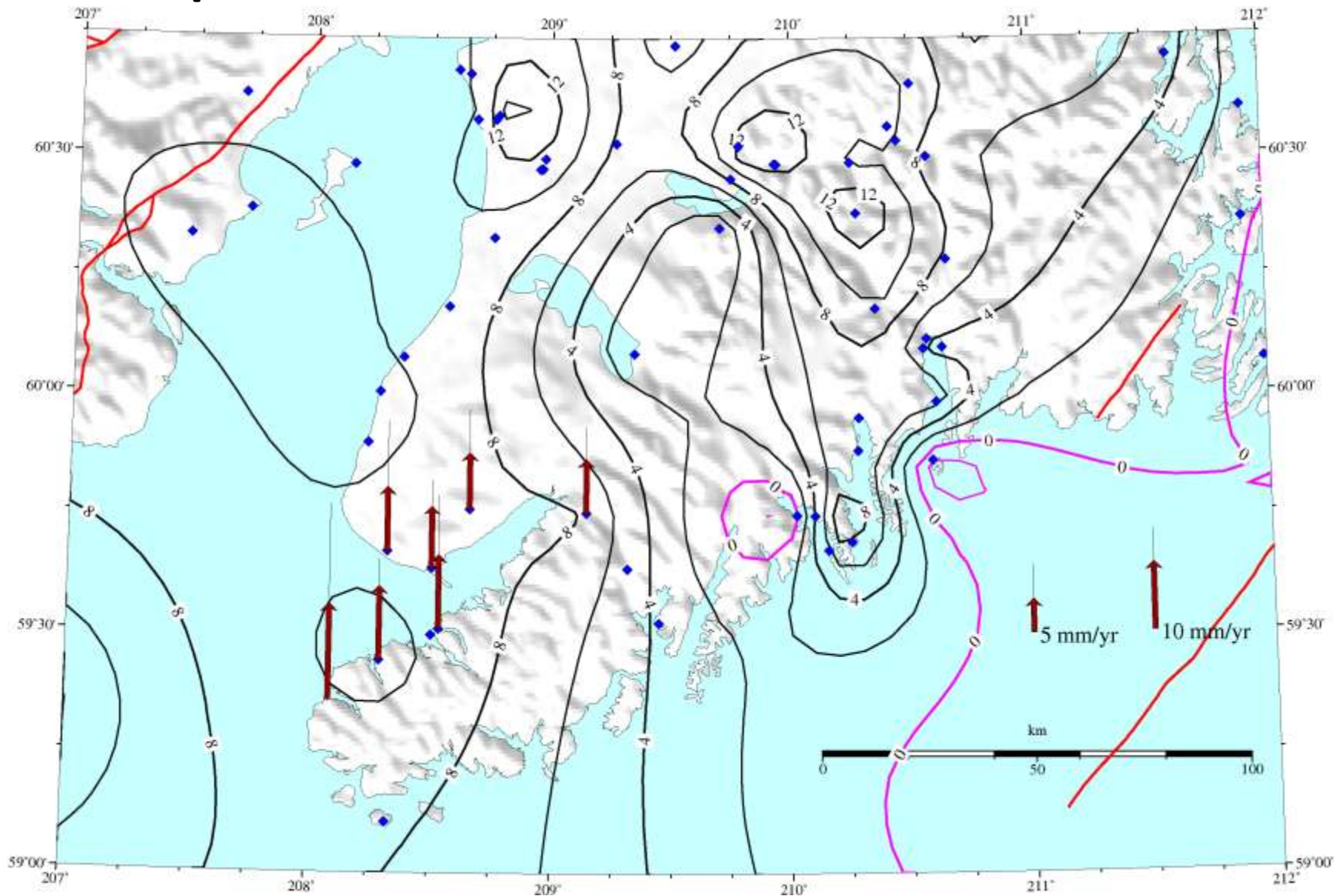
Relative Sea Level



Relative Sea Level

- RSL = *level of the sea* minus *level of the land*
- Both terms can vary regionally
 - Level of the sea: thermal expansion, addition of water, **gravity field**, oceanography
 - Level of the land: **tectonics**, **post-glacial rebound**, **changes in water/ice loading**, **compaction of sediments**, etc.
- In Alaska, land level changes are, in general, more rapid than sea level changes.

Uplift Rates – Kenai Peninsula



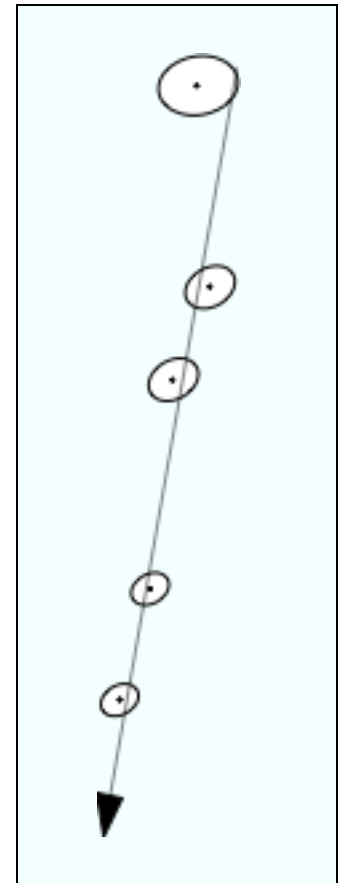
Global Positioning System



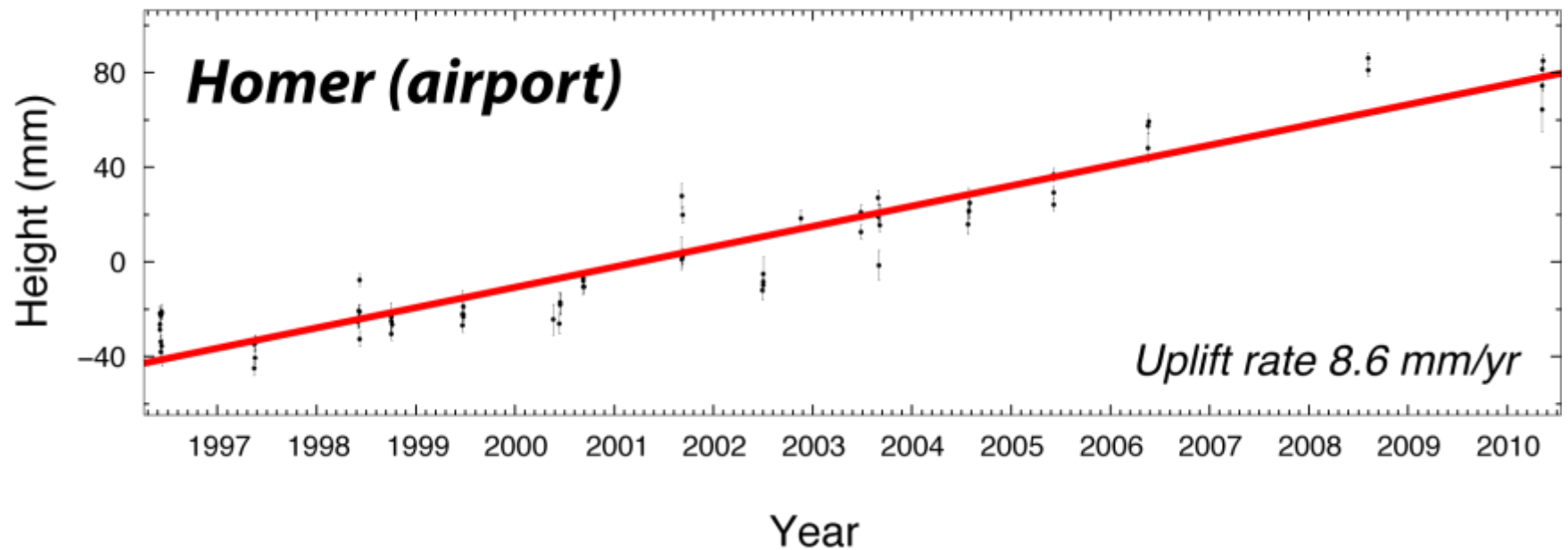
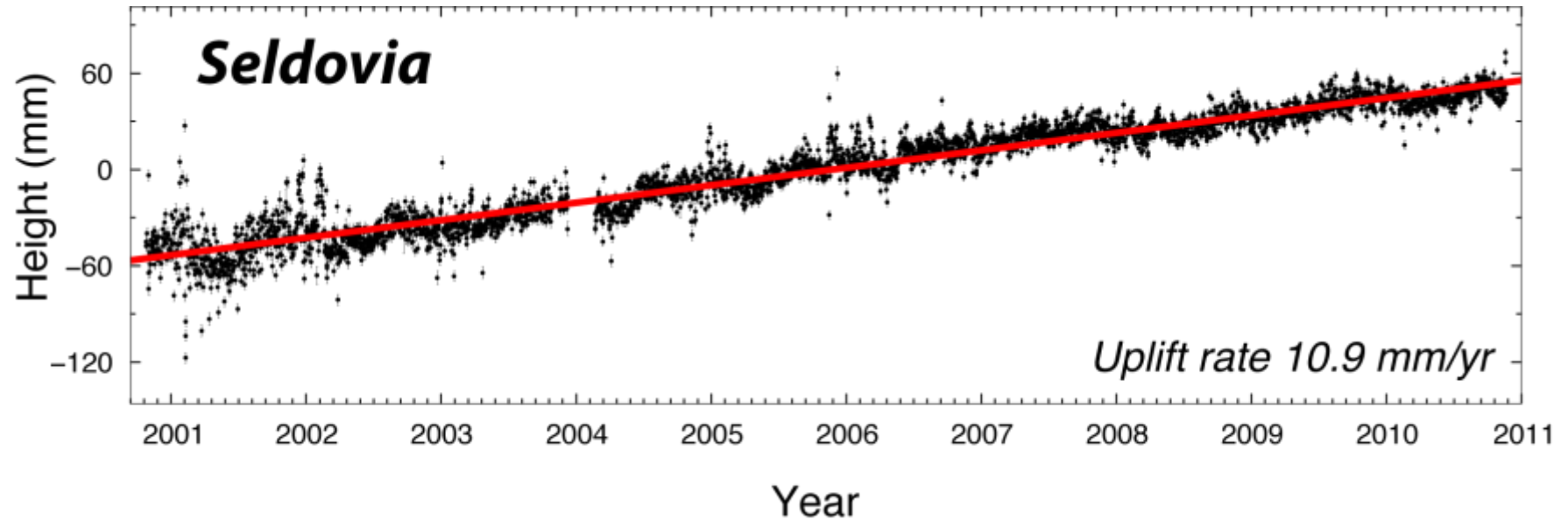
- Portable Surveying Equipment
- Precision of a few millimeters in 3D
- Repeated surveys measure motion of sites

Measuring the Crust

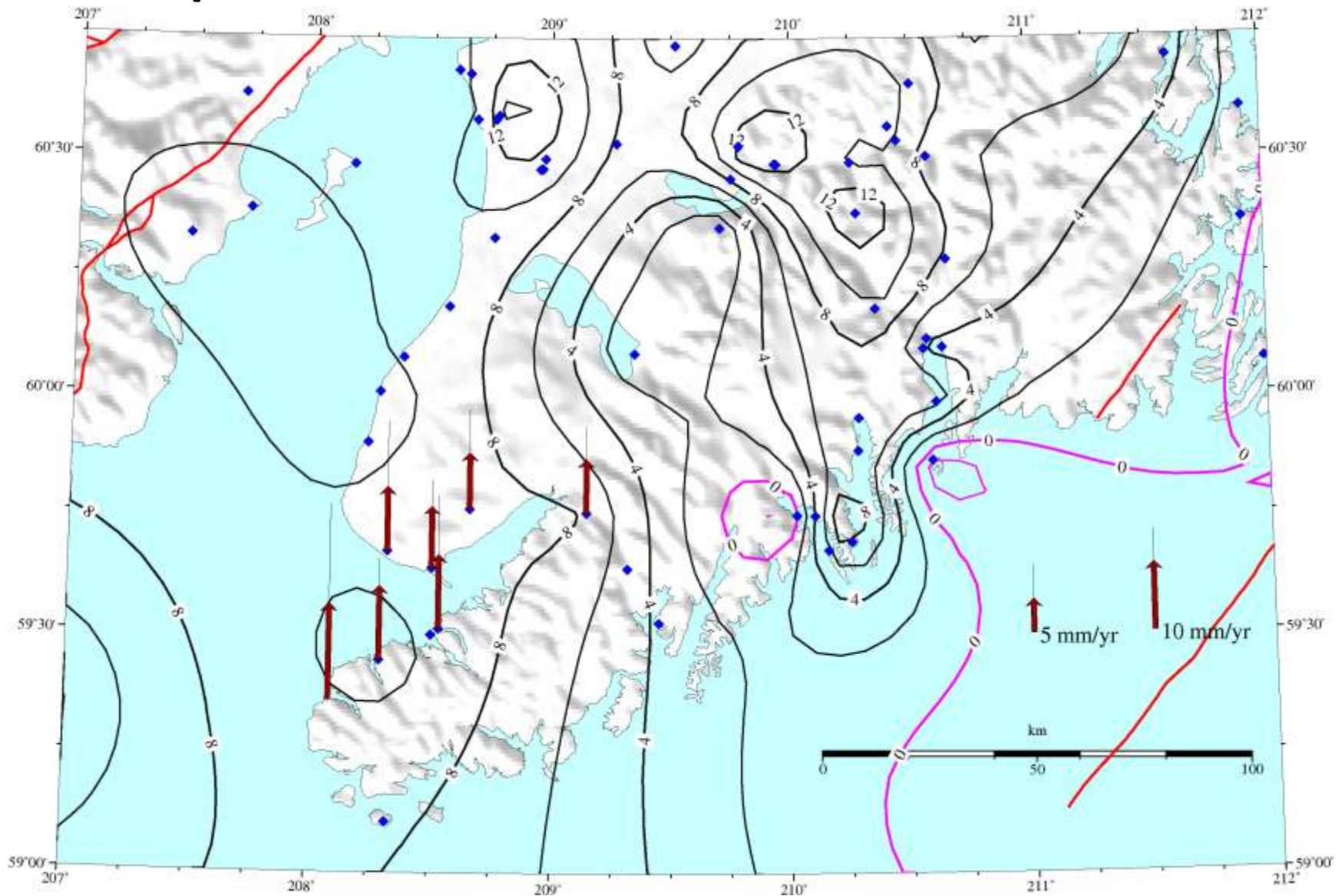
- GPS surveys repeated over time
- Series of positions records the motion of a point fixed to the crust
 - Plate motion
 - Deformation
 - Measurement noise
- Three Dimensions!



Vertical Time Series



Uplift Rates – Kenai Peninsula

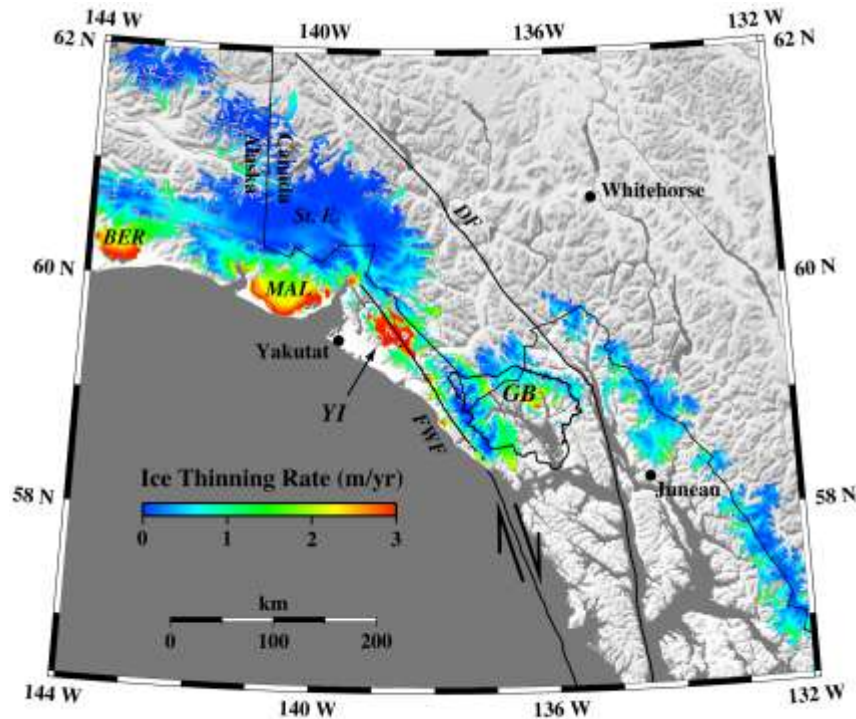


Main Causes of Uplift Locally

- Post-Glacial Rebound/Glacial Isostatic Adjustment
 - Uplift of the land caused by removal of ice load
 - Time-delayed (viscous) response due to past load changes
 - Instant (elastic) response due to ongoing load changes
- Tectonic effects
 - Uplift due to post-seismic deformation after 1964 earthquake
 - Deformation due to plate coupling at subduction zone

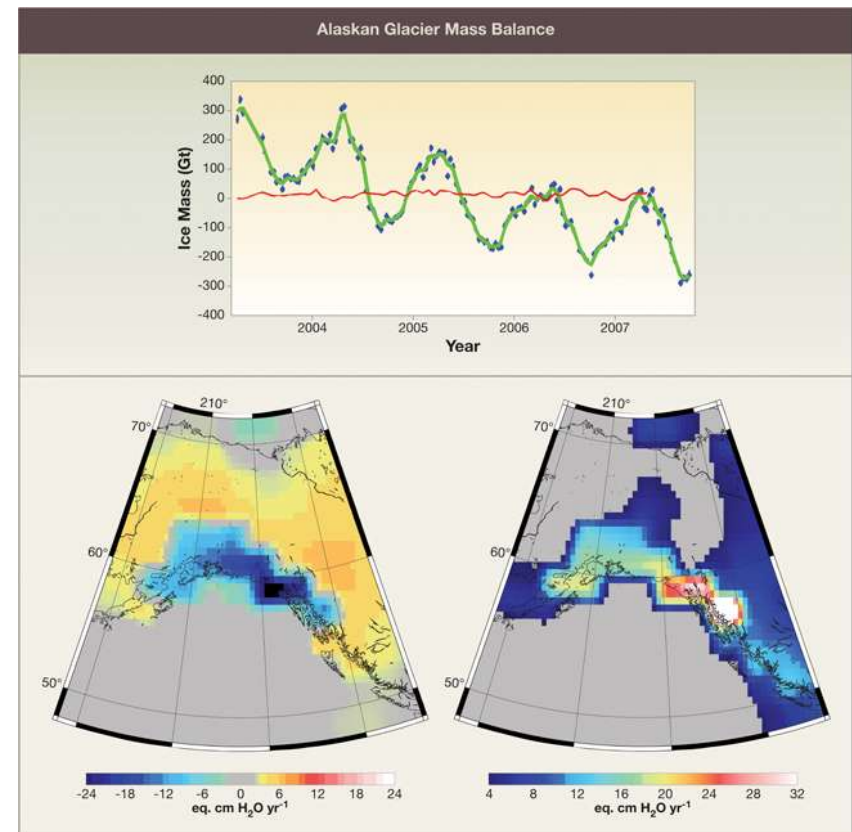
Post-Glacial Rebound – Melting Ice

From repeat glacial altimetry



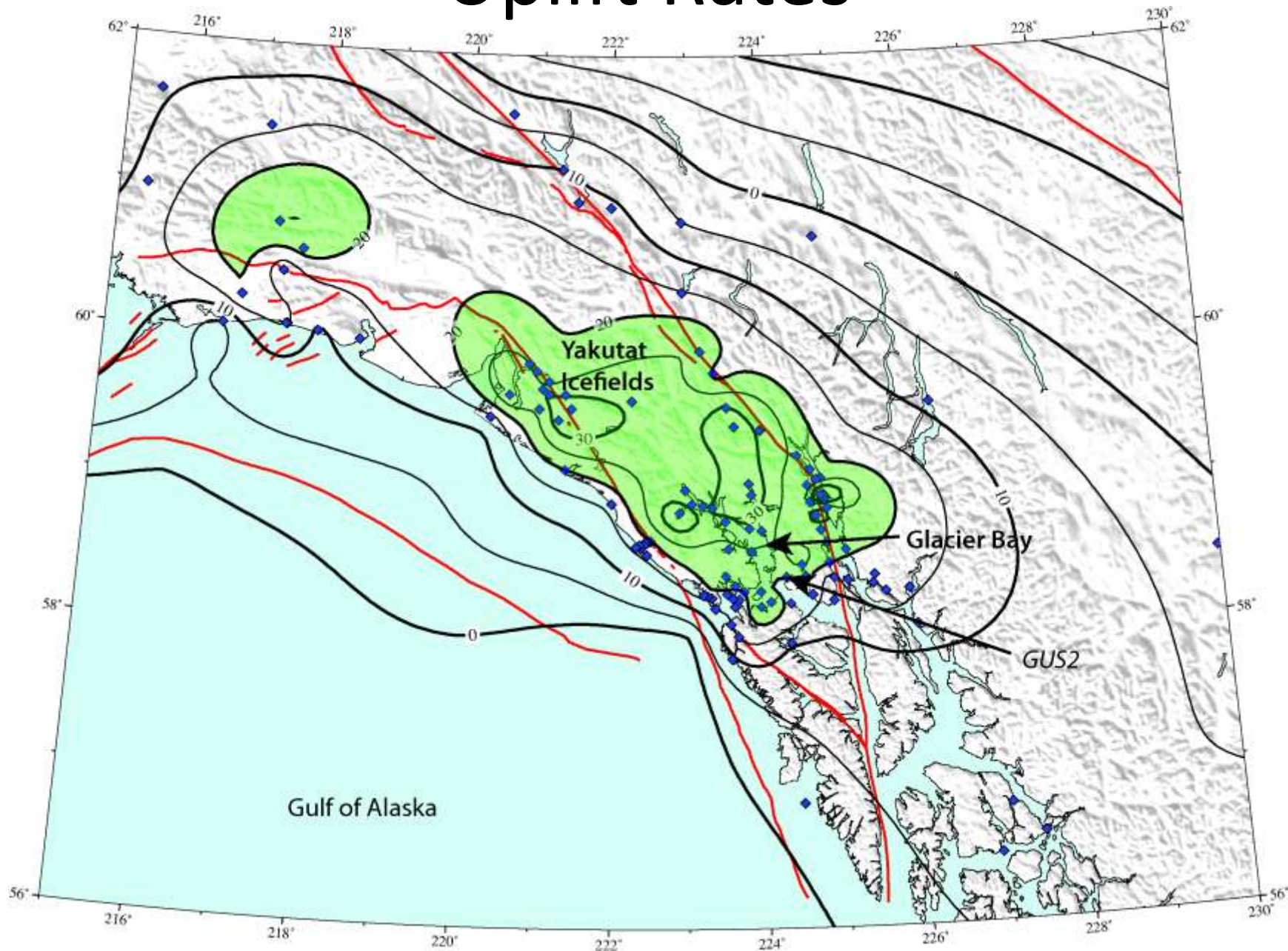
Arendt et al. (2002)

From geoid changes (GRACE)

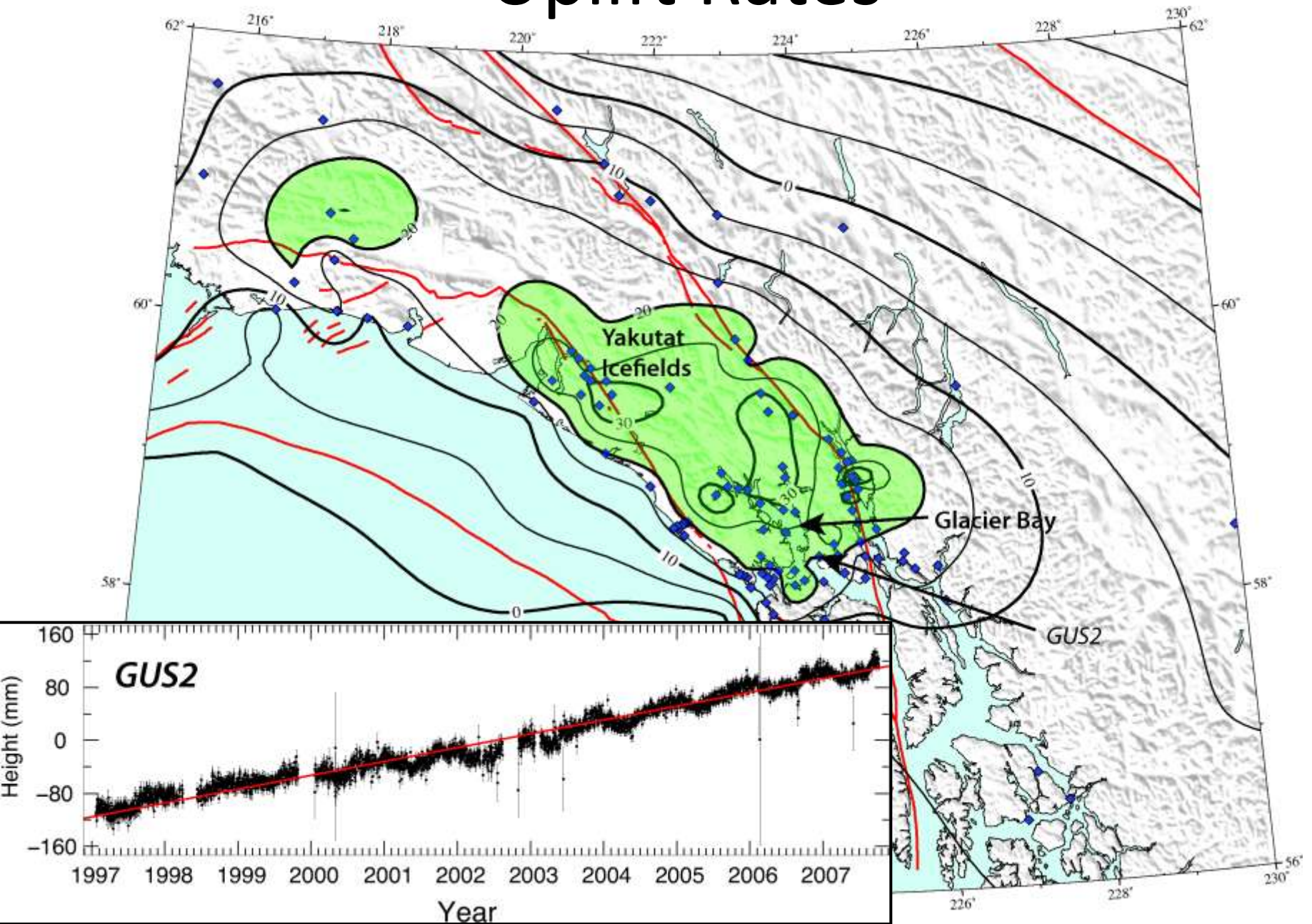


Luthcke et al. (2008)

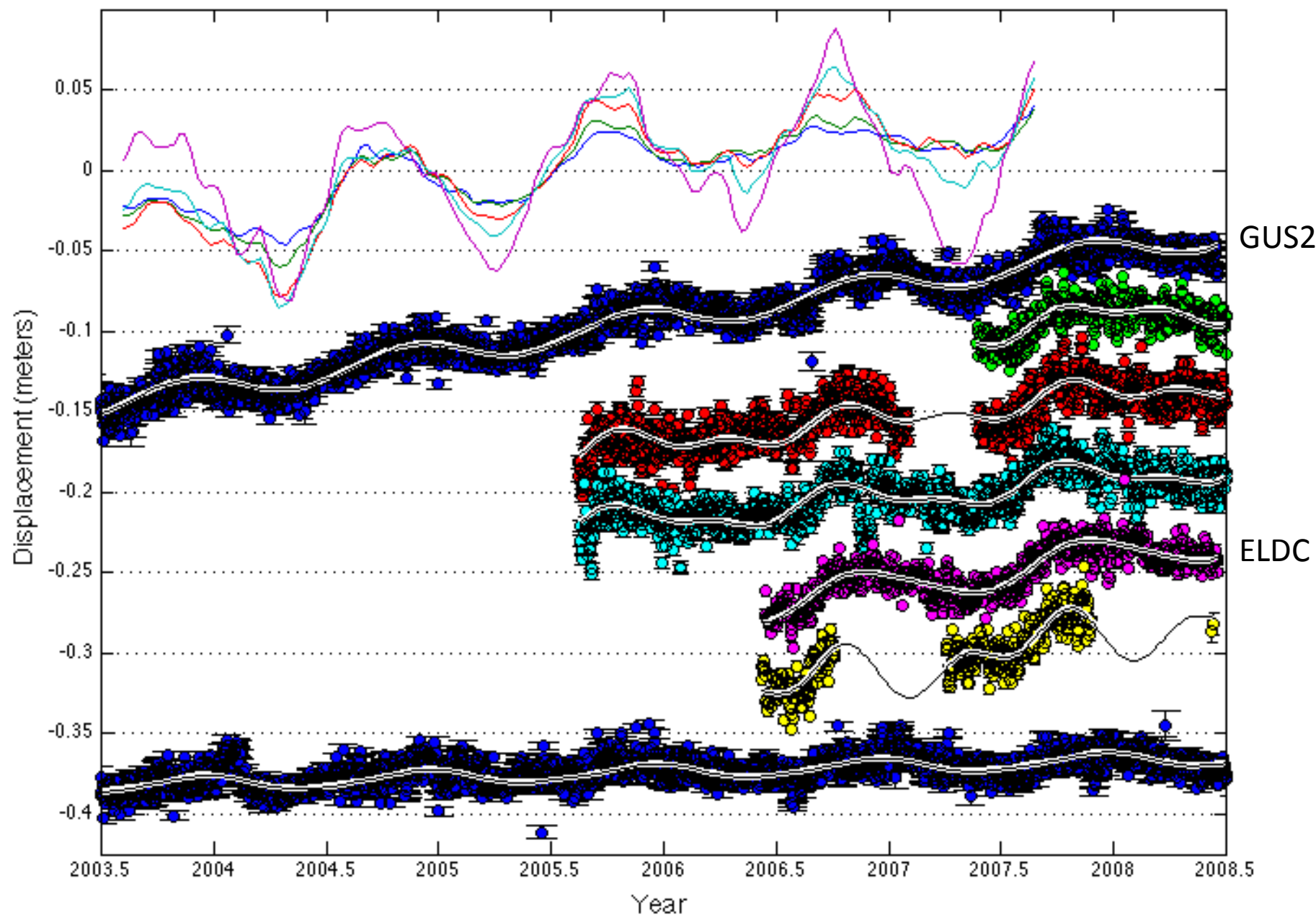
Uplift Rates



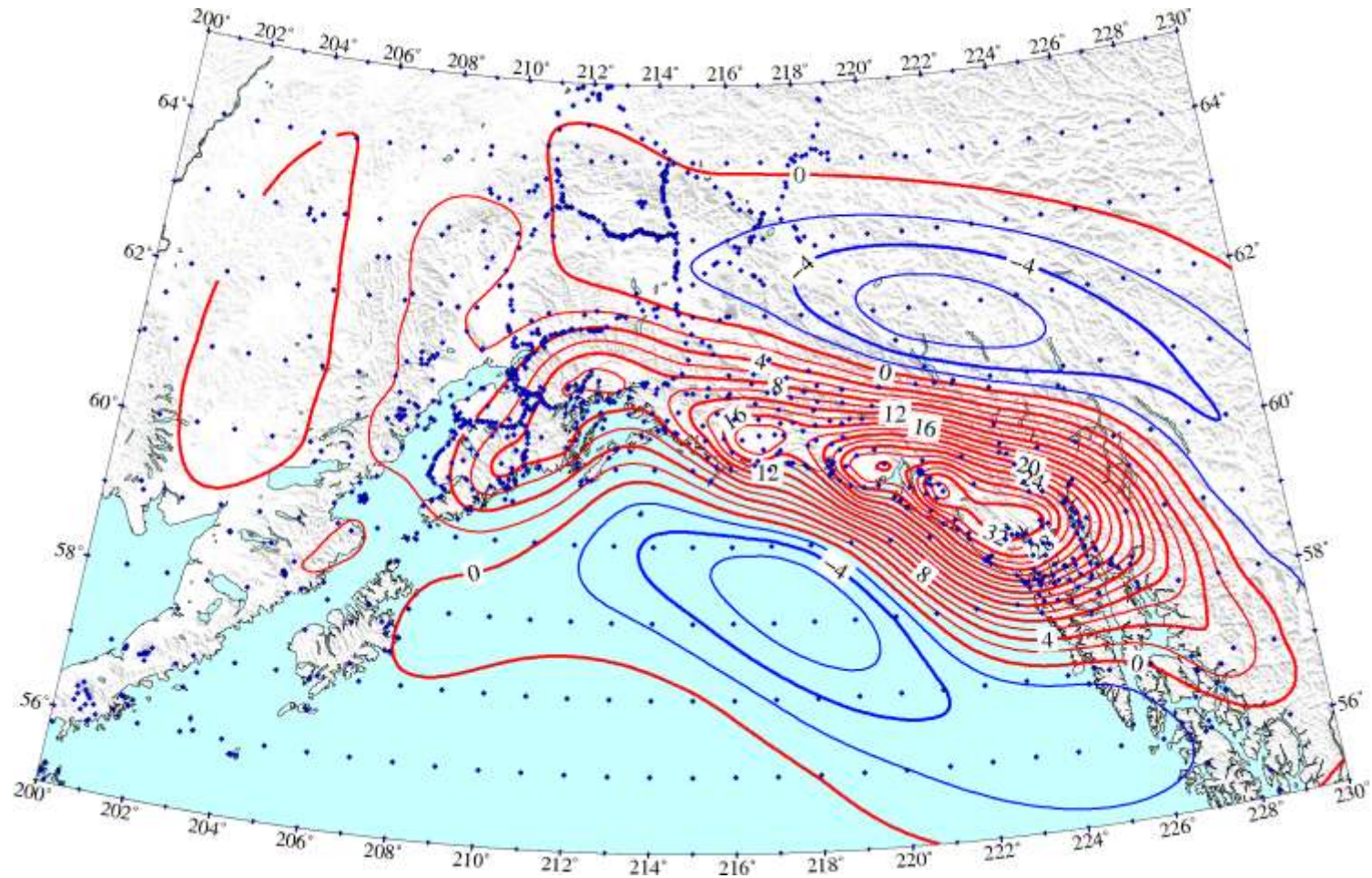
Uplift Rates



GPS and GRACE variations

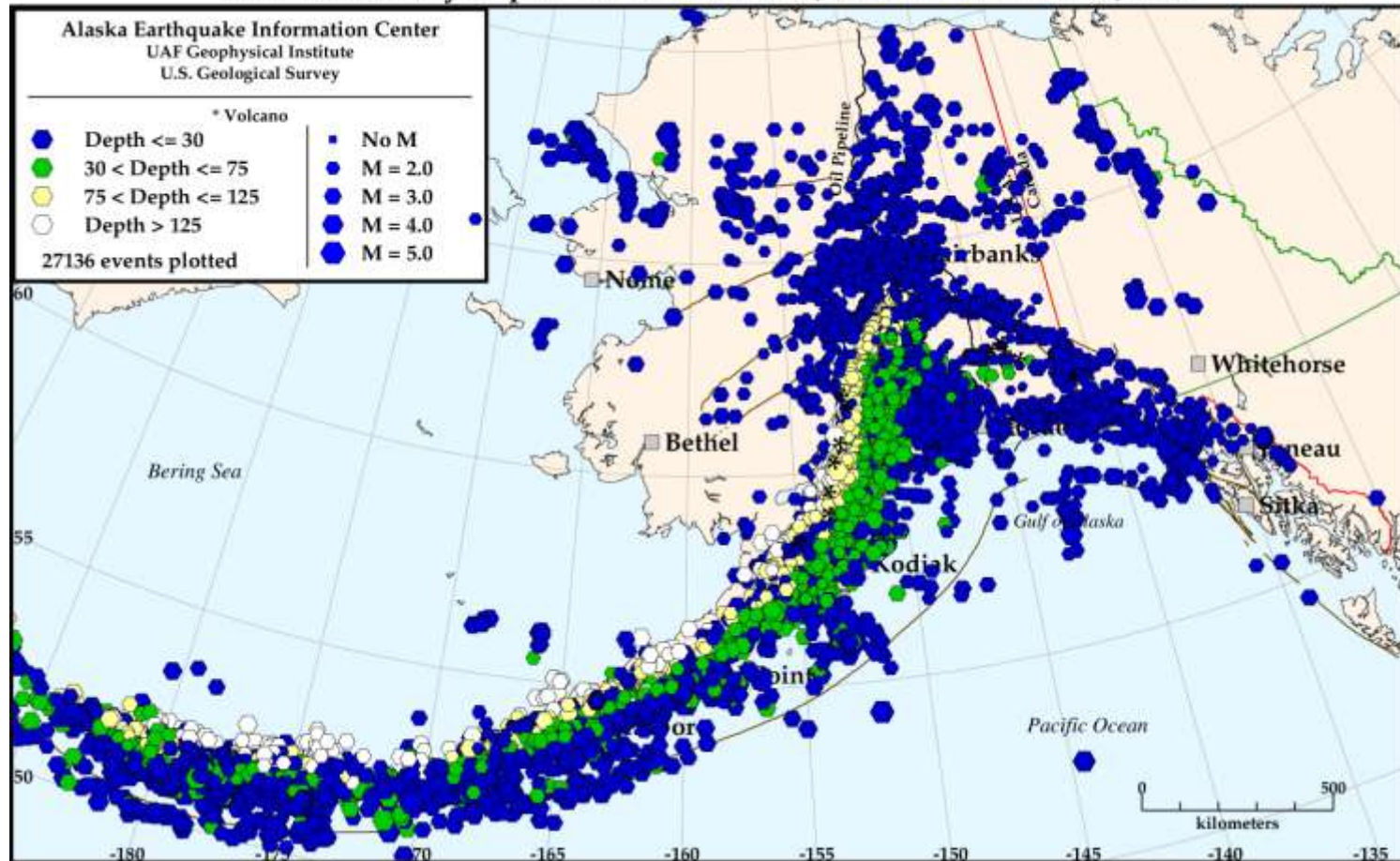


Predicted Uplift/Subsidence Rates

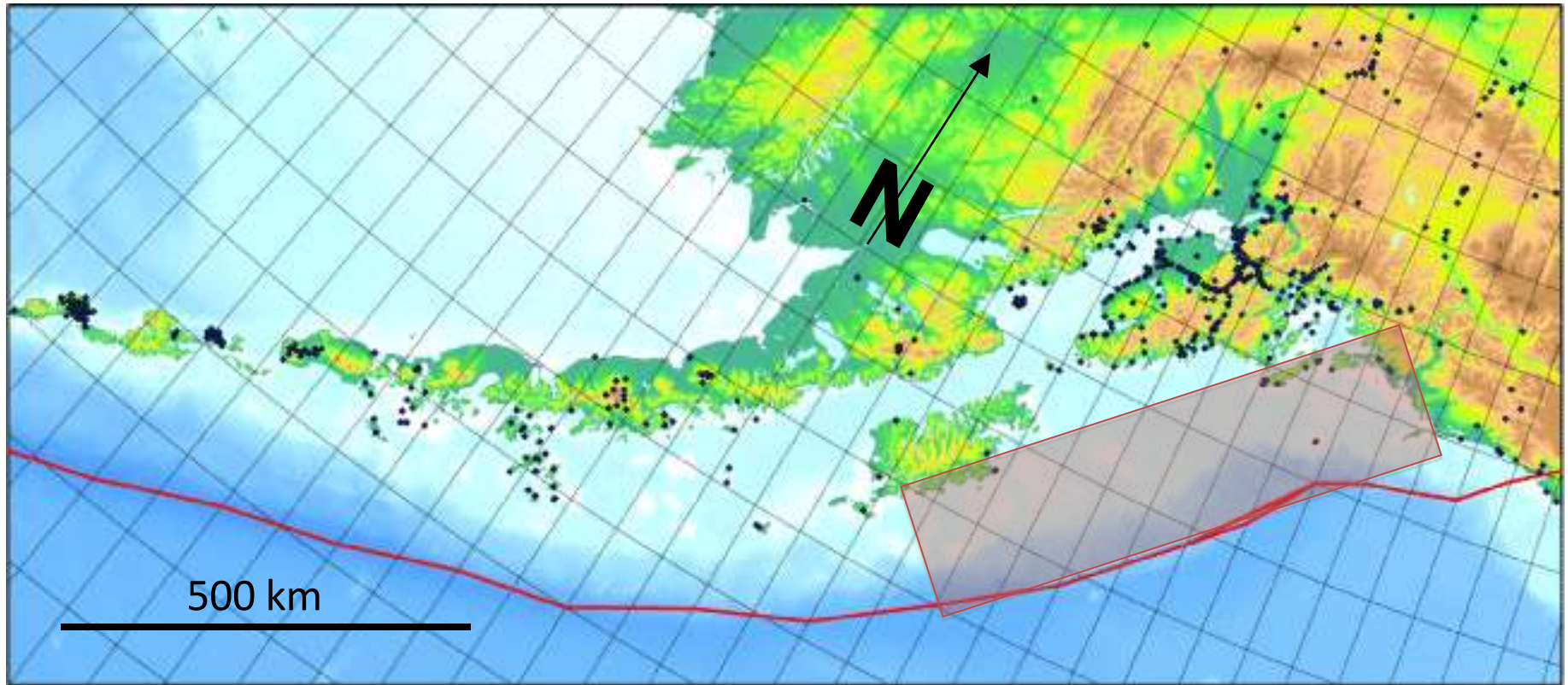


One Year's Seismicity

AEIC Seismicity Report for December 01, 2007 - November 15, 2008

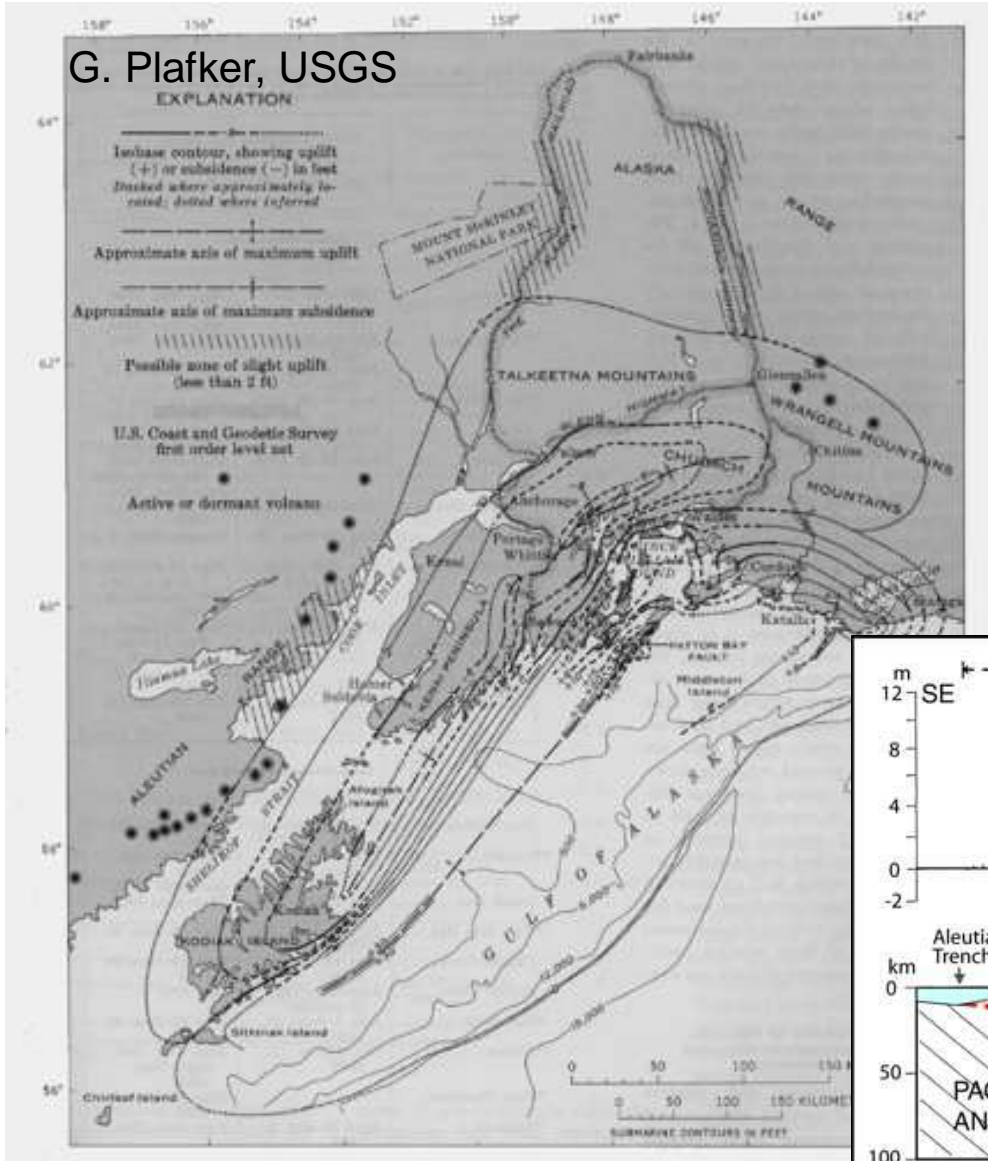


Tectonic and Earthquake Effects

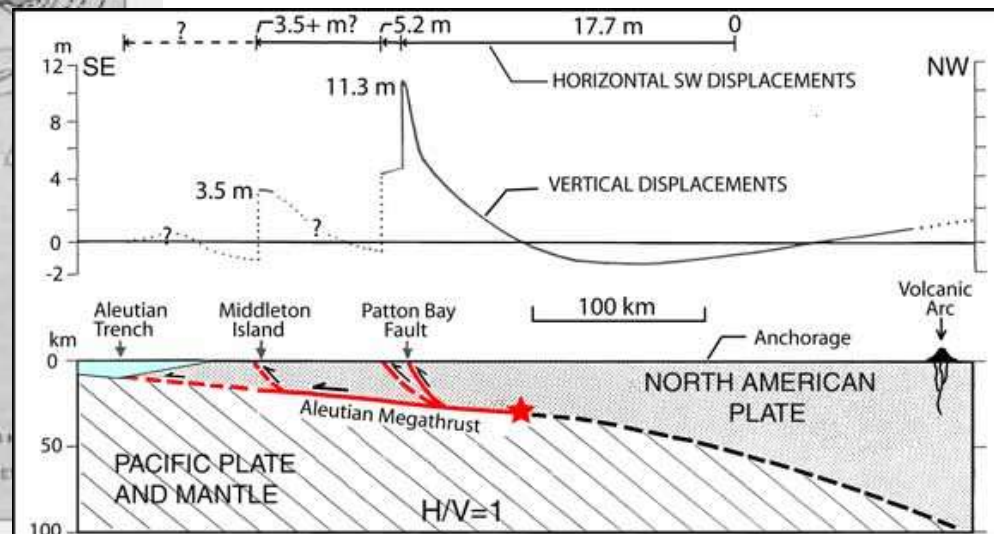


Subsidence in 1964

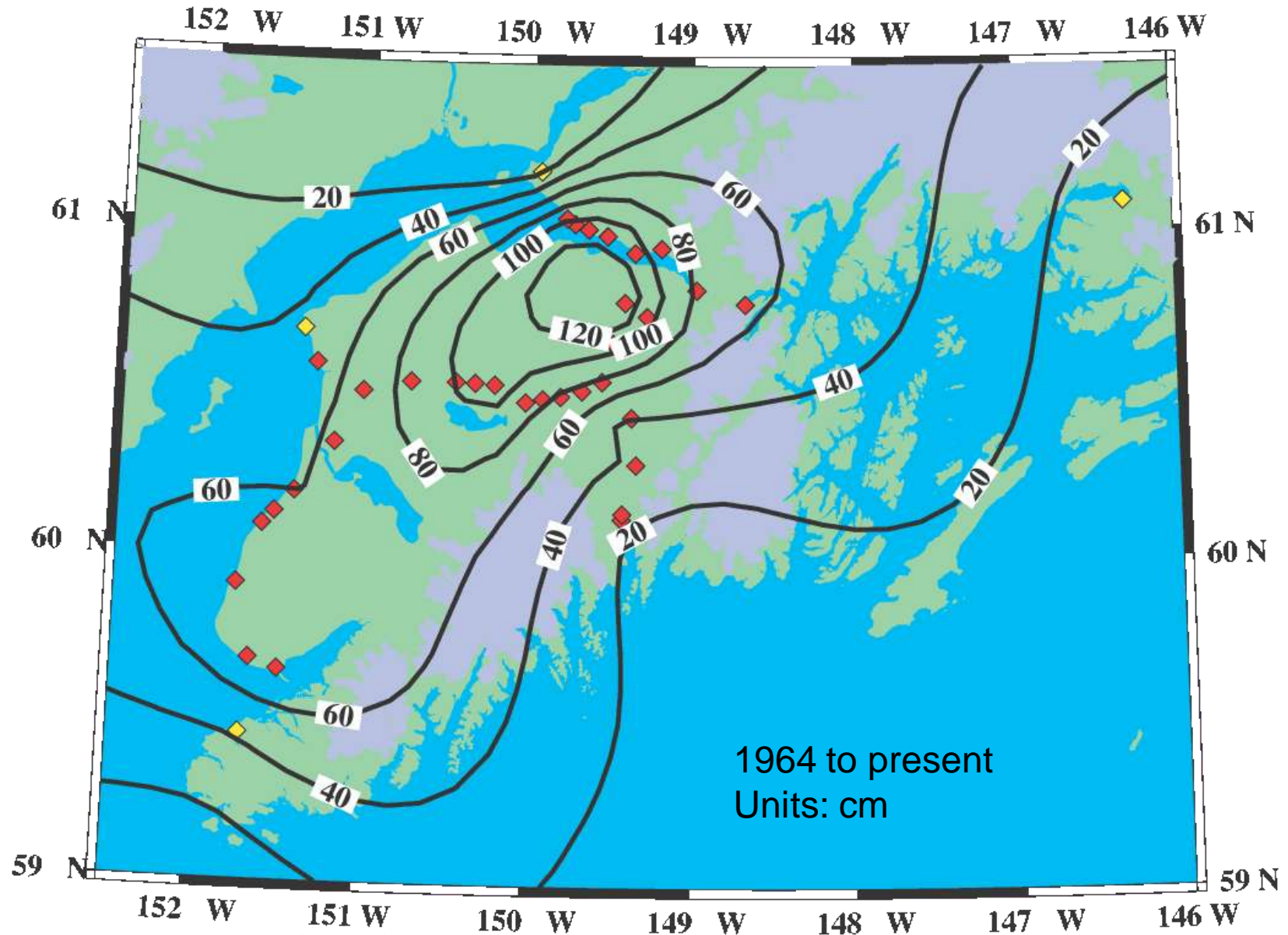
G. Plafker, USGS



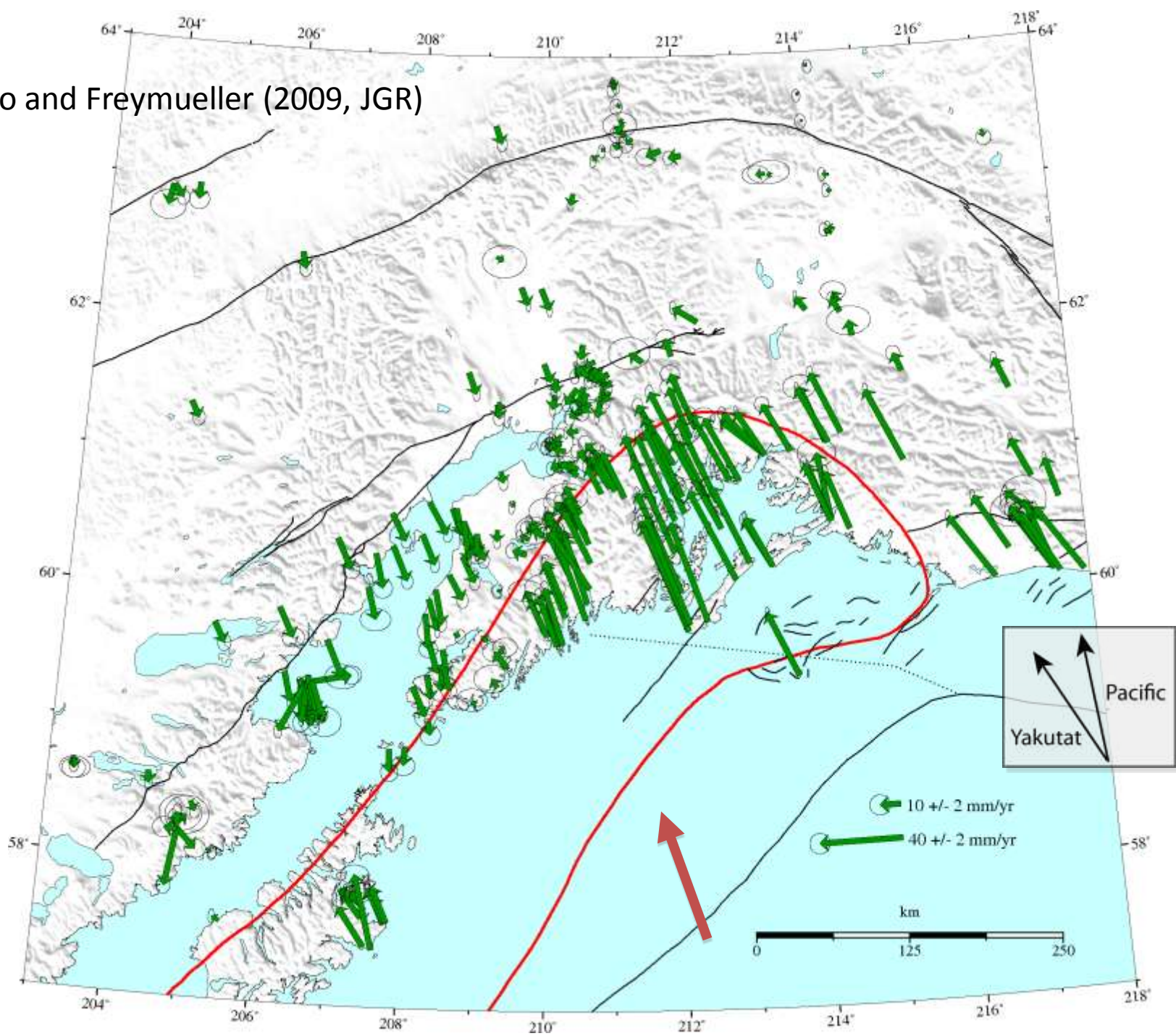
- Kachemak Bay lay near axis of maximum subsidence in 1964
- About 1 meter subsidence
- Cyclic pattern
 - Land here uplifts between earthquakes and subsides during earthquakes

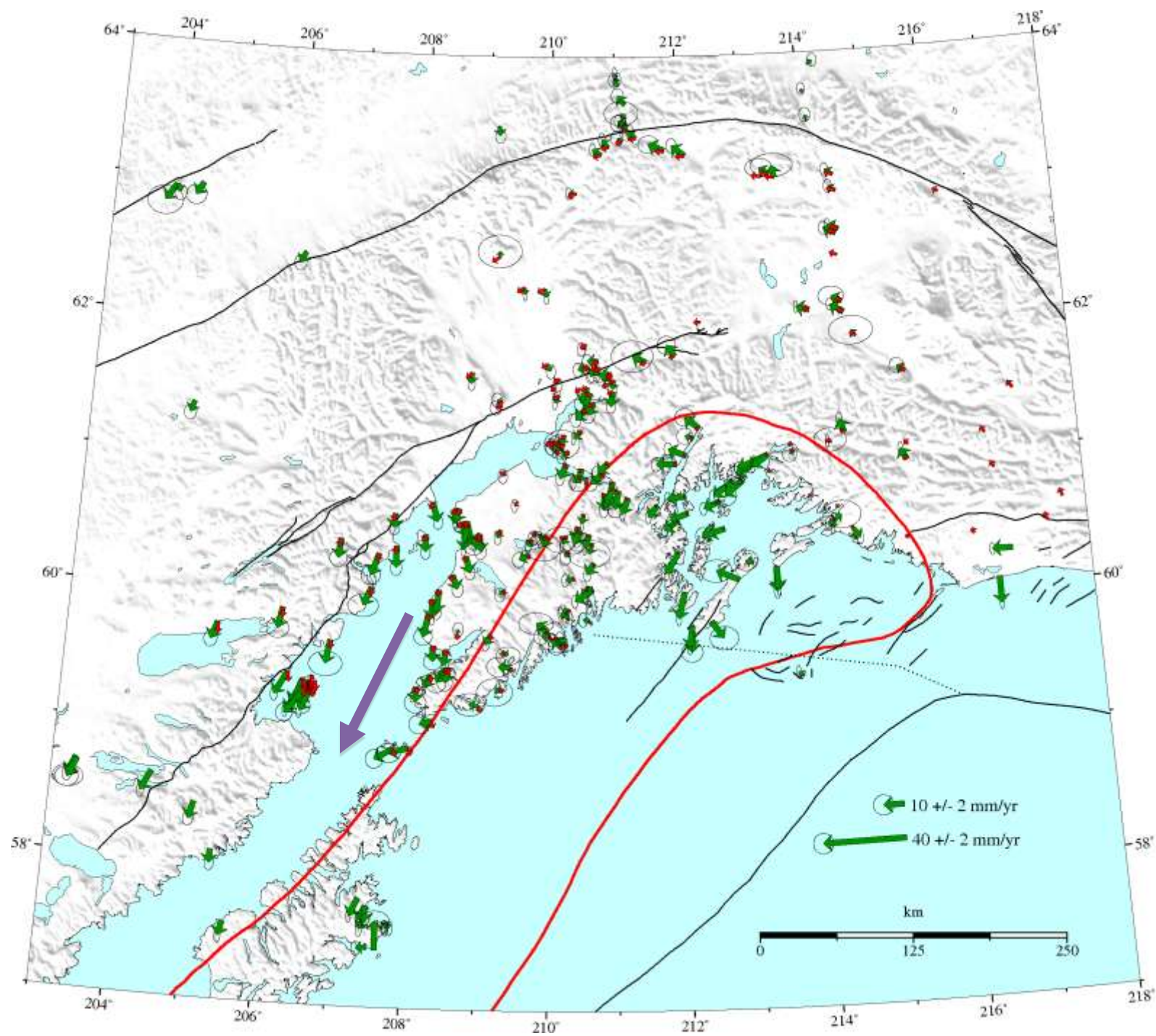


Post-1964 Postseismic Uplift



Suito and Freymueller (2009, JGR)





Knowns and Unknowns

- Can we measure and/or predict vertical motions precisely enough
 - How far into the future?
 - Available models explain the horizontal observations well. Do they explain vertical? Or is there something missing?
- Bedrock vs. soft sediments
 - Sediments self-compact – how fast?
 - Silt-rich tidal flats vs. “old” sediments in town

Project Goals

- Measure vertical motions across Kachemak Bay more precisely
 - Precise enough for useful relative sea level predictions
 - Verify that uplift rates are not uniform in region
- Extend models that work for horizontal motions to explain vertical
- Project future relative sea level change locally
- Combine with response of ecosystems

Continuous GPS sites (CORS)



